

Introduction of the Hungarian Detailed Soil Hydrophysical Database (MARTHA) and its use to test external pedotransfer functions

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The lack of sufficient data of soil hydraulic properties often limits the application of different simulation models (e.g. those calculating crop growth, nutrients and pollutants dynamics, CO₂ sequestration, soil organic matter dynamics). Current methods for direct measurements of soil hydraulic properties are complex, time-consuming and costly. Consequently, there has long been interest in methods for estimating soil hydrological parameters from commonly available (more easily or routinely measured) soil parameters, such as particle-size distribution, bulk density and organic matter content (e.g. COSBY et al., 1984; AHUJA et al., 1985; RAJKAI, 1988; VERECKEN et al., 1989, VAN GENUCHTEN & LEIJ, 1992; RAJKAI et al., 2004). Additional laboratory soil parameters are rarely used (WÖSTEN et al., 2001). Approaches used for generating soil information from existing soil parameters are called pedotransfer functions (PTFs). The term was introduced by BOUMA & VAN LANEN (1987), and became of common use after BOUMA (1989).

In addition to easily measured soil properties other information, as field topographic parameters (PACHEPSKY et al., 2001; RAWLS & PACHEPSKY, 2002) and the combination of physical data and terrain attributes (ROMANO & PALLADINO, 2002), were also used to predict soil hydraulic properties. ANDERSON and BOUMA (1973) and BOUMA et al. (1979) predicted saturated hydraulic conductivity (K_{sat}) values from morphometric soil data. MCKEAGUE et al. (1982) related measured K_{sat} values to soil structure, porosity, biopores, soil texture, consistency and density. They used their results for creating K_{sat} classes for soil horizons. BOUMA (1989) or CRESSWELL et al. (1999) used functional morphologic descriptors (e.g. areal porosity, structure grade and field estimated aggregate stability) to describe the morphological characteristics of soils, because it is difficult to analyse conventional morphological data (used for soil classification and mapping) statistically. The taxonomical classes (soil type or subtype) can also integrate a lot of unknown or less

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known physical, chemical, mineralogical and morphological properties of soils (TÓTH et al., 2008). In the case of brown forest soils the soil water retention characteristics were efficiently estimated using the relating grouped mean data and the soil subtype, texture and humus data codes of the soil maps (MAKÓ et al., 2005). TÓTH et al. (2006) sufficiently predicted the water retention and the hydraulic conductivity values at a selected case-study area according to the category-type data of the soil maps.

Recently a large number of estimation procedures have been developed, using different methods, based on multiple linear regression (e.g. GUPTA & LARSON, 1979; VERECKEN et al., 1989; WÖSTEN et al., 1999; RAJKAI et al., 1999), on physical-empirical approaches (e.g. ARYA & PARIS, 1981), on neural networks (e.g. PACHEPSKY et al., 1995; SCHAAP et al., 1998; MINASNY & MCBRATNEY, 2002), or on classification and regression trees (CART) (PACHEPSKY et al., 2006).

The availability of a thoroughly checked hydrophysical soil database is a prerequisite for using the above-mentioned different data mining tools and for developing pedotransfer functions. A number of hydrophysical soil databases have been constructed in the last two decades worldwide. UNSODA (Unsaturated Soil Hydraulic Database) is an international database of unsaturated soil hydraulic properties, including information on water retention, hydraulic conductivity, soil water diffusivity and basic soil properties. It contains approximately 800 data sets (LEIJ et al., 1996). The HYPRES (Hydraulic Properties of European Soils) database – developed by 20 institutions from 12 European countries (WÖSTEN et al., 1999) – contains measured soil hydraulic characteristics for 5521 soil horizons. The International Soil Reference and Information Centre (ISRIC) prepared a uniform soil data set for the development of pedotransfer functions. The necessary chemical and physical soil data have been derived from ISRIC's Soil Information System (ISIS) and the CD-ROM of the Natural Resources Conservation Service (USDA-NRCS). The name of this data set is IGBT-DIS soil database and contains data for 131,472 samples, originating from 20,920 profiles (TEMPEL et al., 1996). HODNETT and TOMASELLA (2002) selected data from the IGBT-DIS soil database, which originated from tropical regions (771 horizons from 21 tropical countries) to develop pedotransfer functions for tropical soils. Using the Swedish soil physical database (2025 soil layers, about 300 soil profiles) a three-parameter van Genuchten-type model was constructed by RAJKAI et al. (1996) to describe the water retention characteristic data of Swedish soils. KÄTTERER et al. (2005) developed pedotransfer functions to estimate plant available water and bulk density from a database of arable soils in Sweden.

In Hungary two databases are available for developing site-specific or national pedotransfer functions. One of these is the dataset of the Research Institute for Soil Science and Agricultural Chemistry of the Hungarian Academy of Science (RISSAC), which contains information about 270 soil samples, mainly from the Great Hungarian Plain. The pedotransfer functions developed on this dataset (e.g. RAJKAI, 1988, RAJKAI et al., 1999) can successfully be applied for Hungarian chernozem soils. The other hydrophysical dataset is the Unsaturated Soil Hydraulic Database of Hungary (HUNSODA), including data of about 840 soil samples and

soil water retention characteristics of 576 soil horizons (NEMES, 2002). Both of these databases are very useful, their disadvantage, however, is that they only provide information about narrow groups of arable soils.

In the framework of Grant No. T048302 provided by the National Scientific Research Fund (OTKA) we had the opportunity to develop the Hungarian Detailed Soil Hydrophysical Database (MARTHA) with the collaboration of the County Offices of the Hungarian Plant and Soil Protection Service.

The objective of the presented work was to introduce the newly developed Hungarian MARTHA database (ver2.0), then to test two common pedotransfer functions on the database and to evaluate the accuracy of the predictions.

Materials and Methods

The description of MARTHA ver2.0 database

Our aim was to collect all of the measured soil hydrophysical data available in Hungary and to harmonize them into a uniform database, called MARTHA (acronym of the Hungarian name of the database), which is the Hungarian Detailed Soil Hydrophysical Database. As data on agricultural areas were received from all over the country, this database is representative for Hungarian soils being under cultivation. It was applied on a database server of SQL platform (Firebird 2.0). The selected program language was Delphi. To visualize the locations of the soil profiles GoogleMap connection was used. The recent version is the MARTHA ver 2.0.

The MARTHA ver2.0 database includes the existing smaller datasets: the above-mentioned dataset of HUNSODA (750 horizons) and the data of the Hungarian Soil Information and Monitoring System (TIM) (4647 horizons) (VÁRALLYAY et al., 2009). Further to these basic datasets, the second main additional data source is that from the Plant and Soil Protection Services of the Hungarian Counties, which produce various purpose soil assessments (e.g. for irrigation planning) and collect data for these needs (9608 horizons). The first period of data collection has closed. The MARTHA ver 2.0 database currently contains the soil physical, chemical data of 15,005 soil horizons belonging to 3,937 soil profiles. Fig. 1 outlines the location of the MARTHA's soil profiles on the topographical map of Hungary.

The language of the program is Hungarian; the English version will be available soon. The management software is executed with the help of menus. The user can select the proper soil sample according to the identifiers of the soil profiles and the number of their horizons. The general, chemical and physical parameters of the soil profiles can be reached from the overlapping sheets. The *General parameters* sheet contains basic information about the soil profile (identifier; origin of the sample; name of the county where the soil profile is located; EOVS coordinates; GPS coordinates; soil type and subtype); the selected soil profile's picture and location on the map (with Google Map connection); horizons of the selected soil profile (name and depth of the horizon). The *Chemical parameters* sheet stores data about the pH(H₂O); pH(KCl); acidity values (y_1 , y_2); calcium carbonate; salt content; exchangeable Na; sum of Na, K, Ca, Mg cations; CEC and organic matter content.

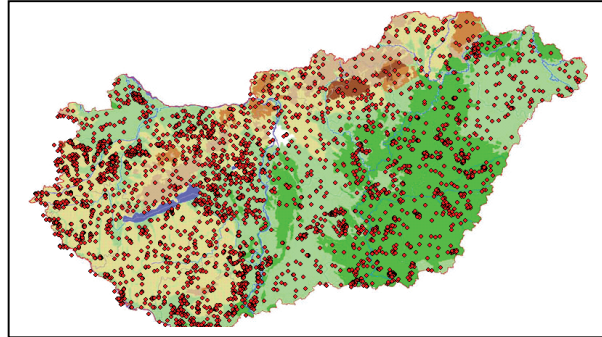


Fig. 1

Locations of soil profiles in the MARTHA ver2.0 database (visualized by the software)

The *Physical parameters* sheet holds the following parameters: soil water retention characteristics (soil water contents at water potentials of -1, -2.5, -10.0, -32.6, -100, -200, -316, -2512, -15850 and -1584893 hPa); particle size distribution (0.25–2; 0.05–0.25; 0.02–0.05; 0.01–0.02; 0.005–0.01; 0.002–0.005; < 0.002 mm); bulk density; specific density; plasticity limit (according to Arany); hygroscopic water content; hydraulic conductivity. Soil properties were measured in accordance with the Hungarian standards (BUZÁS, 1993). Table 1 shows the absolute and relative availability of all *sample* attributes in the database.

Testing pedotransfer functions

One of the potential applications of MARTHA ver2.0 is to develop and apply methods for estimating water retention characteristics from more easily measured data. To establish this work the reliability of two conventional pedotransfer functions (PTFs) was studied.

The Hungarian widely used point estimation PTFs (RAJKAI, 1988; RAJKAI & VÁRALLYAY, 1989) were tested first. Secondly the continuous PTFs of WÖSTEN et al. (1999), developed on the HYPRES European database, were applied.

As all soil parameters selected for the two prediction methods were not necessarily measured for all soil samples (Table 1), there were a number of gaps in the MARTHA ver2.0 database. Therefore the database was checked and reduced to 7524 horizons. This condensed database was used in the further statistical examination.

The quality of the estimated retention values was evaluated using the determination coefficient, R^2 coefficient (for correlation between the measured and the estimated soil water contents at several pF values) and RMSR (Root mean squared residual) (NEMES et al., 2003). Following RAJKAI et al. (2004), the goodness of the predicted water retention was also evaluated. The prediction was considered 'good' if the RMSR was less than 2.5%. The estimation efficiency (EE) of the different estimation procedures was defined as the percentage of 'good' predicted soils.

Table 1
Characterization of the MARTHA ver2.0 database

Attribute	Count	%
CLASS_HUN (soil classification – HUN – national)	14 748	98.3
CLASS_WRB (soil classification – WRB)	14 748	98.3
CLASS_SOILTAX (soil classification – Soil Taxonomy)	14 748	98.3
HOR_N (horizon number)	15 005	100.0
HOR_DES (horizon designation)	14 995	99.9
TOP_D (depth of the top of the sample, cm)	14 948	99.6
BOT_D (depth of the bottom of the sample, cm)	14 897	99.3
OM (weight% of organic matter – Tyurin method)	13 388	89.2
CAR (weight% of free CaCO ₃)	11 455	76.3
KA (plasticity limit according to Arany)	14 696	97.9
PH_H2O (pH in a 1:2.5 soil–water suspension)	14 770	98.4
PH_KCL (pH in a 1:2.5 soil–KCl suspension)	7 542	50.3
Y1 (Na ₄ OAc. extractable acidity, in cmol(+) kg ⁻¹)	3 151	21.0
Y2 (KCl extractable acidity, in cmol(+) kg ⁻¹)	979	6.5
SALT (weight% of total salt content)	11 898	79.3
CSNA (exchangeable Na, in cmol(+) kg ⁻¹)	5 712	38.1
CEC (cation exchange capacity of the soil, in cmol(+) kg ⁻¹)	10 313	68.7
S (sum of Na, K, Ca, Mg cations, in cmol(+) kg ⁻¹)	330	2.2
CLAY_USDA_HUN (weight% of particles < 0.002mm)	14 311	95.4
SILT_USDA (weight% of particles 0.05–0.002 mm)	14 322	95.4
SAND_USDA (weight% of particles 2.0–0.05 mm)	14 311	95.4
SILT_HUN (weight% of particles 0.02–0.002 mm)	14 313	95.4
SAND_HUN (weight% of particles 2.0–0.02 mm)	14 310	95.4
FINE_SILT_1 (weight% of particles 0.005–0.002 mm)	14 313	95.4
FINE_SILT_2 (weight% of particles 0.01–0.005 mm)	14 313	95.4
FINE_SILT_3 (weight% of particles 0.02–0.01 mm)	14 313	95.4
COARSE_SILT (weight% of particles 0.05–0.02 mm)	14 313	95.4
FINE_SAND (weight% of particles 0.25–0.1 mm)	14 064	93.7
COARSE_SAND (weight% of particles 1.0–0.5 mm)	14 064	93.7
HY1 (hygroscopic water content (weight%))	5 061	33.7
BD (bulk density – oven dry at 105 °C – in kg·dm ⁻³)	12 629	84.2
DENS (density, in t·m ⁻³)	440	2.9
PF_0 (moisture content at -1 hPa = pF0 (vol%))	12 739	84.9
PF_04 (moisture content at -2.5 hPa = pF0.4 (vol%))	999	6.7
PF_1 (moisture content at -10 hPa = pF1.0 (vol%))	1 849	12.3
PF_15 (moisture content at -32.6 hPa = pF1.5 (vol%))	4 036	26.9
PF_2 (moisture content at -100 hPa = pF2.0 (vol%))	7 870	52.4
PF_23 (moisture content at -200 hPa = pF2.3 (vol%))	889	5.9
PF_25 (moisture content at -316 hPa = pF2.5 (vol%))	12 663	84.4
PF_34 (moisture content at -2512 hPa = pF3.4 (vol%))	1 753	11.7
PF_42 (moisture content at -15850 hPa = pF4.2 (vol%))	12 759	85.0
PF_62 (hygroscopic water content (vol%)) (-1584893 hPa = pF6.2)	12 483	83.2
K_SAT (saturated hydraulic conductivity – cm·day ⁻¹)	2 879	19.2

Optimum partitioning of databases with classification trees (SPSS TREES CHAID) was used to find the best grouping of samples according to the RMSR values (SPSS, 2001). It was assumed that water retention is also affected by soil structure and therefore the quality of estimation procedures depends on the soil structural groups. Because of the lack of direct soil structural data in the MARTHA ver2.0 database, the empirical knowledge about the different soil classification categories (soil types and subtypes) was used to predict the characteristic structural parameters of the soils. For this purpose, the grade of structure and the shape of aggregates were described. Four classes were used to characterize the grade: 1. *structureless*, where no observable aggregation occurs, 2. *poorly structured*, when the aggregates are slightly observable, 3. *medium structured*, when aggregates are well formed in the undisturbed soil, but the stability of the aggregates is moderate, 4. *well structured*, when aggregates are distinct in the undisturbed soil and, the stability of the aggregates is high. The shape of aggregates was classified into five classes: 1. *not aggregated*, 2. *granular*, 3. *prismatic*, 4. *blocky*, and 5. *columnar*. These structural parameters were used as grouping variables to define the soil groups with distinctly different RMSR values.

Results and Discussion

The statistical evaluation of the two estimation procedures is given in Table 2. In general, the application of both examined pedotransfer functions (RAJKAI, 1988; WÖSTEN et al., 1999) was not very successful. This is presumably due to the fact that the data set used by Rajkai originated mainly from the Great Hungarian Plain. Thus, it represented only a few Hungarian soil types. On the other hand, the HYPRESS database used by WÖSTEN et al. (1999) came from different European countries, where soil forming conditions may differ significantly from the Hungarian circumstances. These results suggest that it is necessary to develop new PTFs on the basis of the new Hungarian database, which would be representative for a wider range of the Hungarian soils.

In the second part of our analysis the effect of structural differences on the goodness of estimation was examined. The RMSR values calculated for the classi-

Table 2
Goodness of estimation indicators (R^2 ; RMSR; EE) for different types of PTFs

Estimation procedure	R^2 coefficient			RMSR (vol%)	EE (%)
	pF 0	pF 2.5	pF 4.2		
Measured vs. RAJKAI (1988)	0.68	0.53	0.54	4.715	16.6
Measured vs. WÖSTEN et al. (1999)	0.66	0.53	0.54	5.182	9.1

Remark: R^2 = determination coefficient; RMSR = root mean squared residual; EE = estimation efficiency

fication trees using the two different pedotransfer functions are shown in Fig. 2. The grade of aggregation was the best grouping parameter in both cases. A stronger grade decreased the RMSR values (increased the goodness of estimation). Based on the observed effect of the aggregate's grade on the estimation, it can be supposed that the databases, which were used for developing the PTFs contained less structureless soils than the MARTHA ver2.0 database did. The shape class was in the list

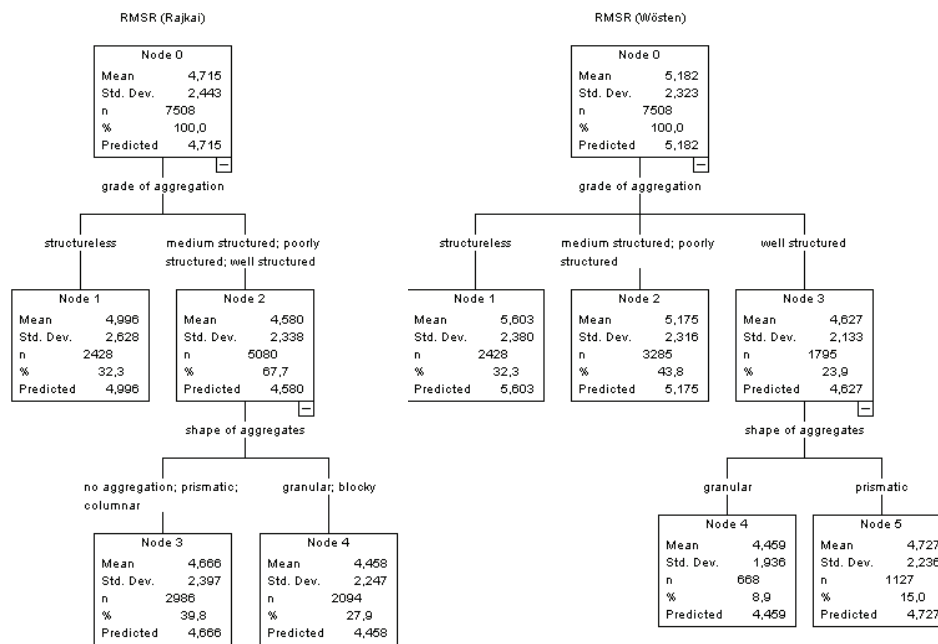


Fig. 2

Classification trees to group soil samples according to root mean squared residual (RMSR) values (vol%)

of grouping parameters only in the case of structured soil samples (Rajkai's PTFs) or in well-structured soil samples (Wösten's prediction). A slight decrease in RMSR values (increase in goodness of estimation) was observed for groups with granular structure. Our hypothesis is that the data sets of both prediction methods represented more soils with granular structure than others. Another reason of this experience may be that salt content, sodium saturation or clay mineral composition cause less estimation errors in the case of soils with granular structure than in soils with e.g. columnar or prismatic structure. Results of the classification tree method show that defining and quantifying the soil structure may partly explain the inaccuracy of soil water retention prediction. Furthermore, the structural data may serve as grouping variable for the further development of class PTFs. As soil forming conditions (represented in the MARTHA database by soil classification units) in Hungary differ to a great extent, structural soil characteristics are greatly variable. Our results

indicate that soil structural properties could serve as important additional information in PTF development, even in the case of their indirect incorporation.

Summary

The Hungarian Detailed Soil Hydrophysical Database, called MARTHA ver2.0 has been developed to collect information on measured soil hydraulic and physical characteristics in Hungary. Recently this is the largest detailed national hydrophysical database, containing controlled information from a total of 15,005 soil horizons.

Two commonly used pedotransfer functions were tested to evaluate the accuracy of the predictions on the MARTHA data set, representative for Hungarian soils. In general, the application of both examined pedotransfer functions (RAJKAI, 1988; WÖSTEN et al., 1999) was not very successful, because these PTFs are representative for other soil groups. The classification tree method was used to evaluate the effect of soil structure on the goodness of estimations. It was found that using the soil structure data the inaccuracies of soil water retention predictions are more explainable and the structure may serve as a grouping variable for the development of class PTFs.

Key words: soil hydrophysical database, pedotransfer functions, water retention

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