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Parametric or non-parametric: analysis of rainfall time series at a Hungarian meteorological station

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Abstract— Parametric methods (linear trend, t-test for slope) for analyzing time series are the simplest methods to get insight to the changes in a variable over time. These methods have a requirement for normal distribution of the population that can be a limit for application. Non-parametric methods are distribution-free methods, and investigators can have a more sophisticated view to the variable tendencies in time series. 144-year-long time series of precipitation data measured at the meteorological station in Keszthely, Hungary (latitude: 46°44', longitude: 17°14', elevation: 124 m above Baltic sea level) were analyzed by Mann-Kendall trend test for detecting tendencies in the time series. Sen's slope estimator was applied to estimate the slope of the linear changes. In average, 44 mm decline can be shown for 100 years in the annual sum, 29.7 mm and 25.7 mm in the precipitation sum of spring and autumn (in 100 years), respectively. The rainfall sum of winter increased by 15.4 mm. Sums of April, May, and October declined by 10.8 mm, 13 mm, and 20.9 mm, respectively, according to one-tailed Mann-Kendall tests. These results were compared to the previous results of the authors carried out by parametric methods. Results of two-tailed tests of parametric and non-parametric methods are easily comparable. Parametric method (linear trend) proved significant decreasing tendencies for spring, April, and October. Non-parametric Mann-Kendall tests show significant declining tendencies for spring, autumn, and October.

Key-words: Kendall's tau, Mann-Kendall trend test, Sen's slope estimator, precipitation changes, Keszthely, Hungary

1. Introduction

Climate change is one of the problems that mankind should face in the 21st century. According to the last *IPCC* report (2013), human role in the process has no doubt (95% is the probability that human influence has been dominant on the present changes of climate system). Climate change will probably affect all parts of the Earth, and in the center of Europe, the Carpathian Region will be influenced as well. In some cases, the volume and direction of the changes in climate model simulations are uncertain. Hydrological cycle is an element of the climate system that is expected to change, and the signs of these amendments can already be detected. Precipitation in average over mid-latitude land areas of the Northern Hemisphere has increased from 1901 (medium confidence) according to *IPCC AR5* (2013). Heavy precipitation events and increase in intensity and frequency of rainfall are very likely (90% probability) over mid-latitude land masses (*IPCC*, 2013). Precipitation strongly influences the water cycle from local to global scales. Any modification in the amount or distribution of rainfall has significant impact on the water availability, and therefore, water management. Hungary, which occupies the middle of the Carpathian-Basin, has long-experienced significant temporal and spatial variations in precipitation. In recent decades, due to the hardships faced by the state complying with decreasing rainfall amount, studies related to precipitation events are extremely beneficial.

The prediction of the effects of climate change on the Carpathian Region (including Hungary) requires regional climate scenarios with adequate temporal and spatial resolution, capable of translating global phenomena to a local scale. *Bartholy et al.* (2004) applied regional models to estimate the regional effects of climate change in the Balaton Lake–Sió Canal catchment area, using ECHAM/GCM outputs. This catchment area (which also includes Keszthely meteorological station) is one of the most vulnerable regions in Hungary because of its economical and touristic importance and unique treasure of nature. According to *Bartholy et al.* (2005), the amount of precipitation will decline by 25–35% in the summer half-year and by 0–10% in the winter half-year on the Balaton Lake–Sió Canal catchment area. The regional model runs for the Carpathian Basin (RCMs) using the A2 and B2 global emission scenarios of the *IPCC AR4* (2007), expect more than 2.5 and less than 4.8 °C temperature rise for all seasons and both scenarios (*Bartholy et al.*, 2007). A 20–33% decrease in precipitation is predicted for the summer half-year and there is high uncertainty for the rainfall for the winter half-year (*Bartholy et al.*, 2007). The earlier results of the authors harmonizes with their latest projection carried out in the PRUDENCE European Project's model application (*Bartholy et al.*, 2009). These statements were enhanced by *Bartholy et al.* (2008) and by the Hungarian Meteorological Service (*OMSZ*, 2010) according to further regional climate model simulations. *Pongrácz et al.* (2011) and *Kis et al.* (2014) communicated similar results for projection of the future climate in Hungary. *Pongrácz et al.* (2014) project significant increase in drought-related indices in summer by the end of the 21st

century. *Bartholy et al.* (2015) examined precipitation indices and project that frequency of extreme precipitation will increase in Central Europe, except of summer, when decreasing tendency is very likely. For the tendencies of the past, *Szalai et al.*, (2005) stated that the annual precipitation amount decreased by 11% between 1901 and 2004, according to the analysis of the Hungarian Meteorological Service. The biggest decline could be experienced in spring; it was 25% for the above mentioned period. *Bodri* (2004) suggested that slow decrease of precipitation with a noticeable increase in precipitation variability are characteristic for the 20th century. While the northern and western part of Europe gets more precipitation in parallel with the warming tendency, Hungary, similarly to the region of the Mediterranean Sea, gets less rainfall. The water balance has deficit, since the difference between water income and outflow is increasing. Between 1901 and 2009, the highest precipitation decline over the territory of Hungary occurred in the spring, nearly 20% (*Lakatos and Bihari*, 2011). *Bartholy and Pongrácz* (2005, 2007, 2010) examined several precipitation extreme indices and suggested that regional intensity and frequency of extreme precipitation increased in the Carpathian Basin in the second half of the last century, while the total precipitation decreased.

The goal of this study is to analyze the long-term data series of the meteorological measurements of precipitation amount at the meteorological station at Keszthely, Trans-Danubia, Hungary (N 46°44', E 17°14', *Fig. 1*) from the point of view of climate and statistics. Our aim is to analyze the dataset by non-parametric methods and compare the results to our previous findings derived from parametric time series analyses. A non-parametric method, the Mann-Kendall trend test, that was applied in this research, is widespread for analyzing meteorological (such as precipitation sums) and hydrological data. The Sen's slope estimator was applied as non-parametric method for determining the slope of the tendencies.

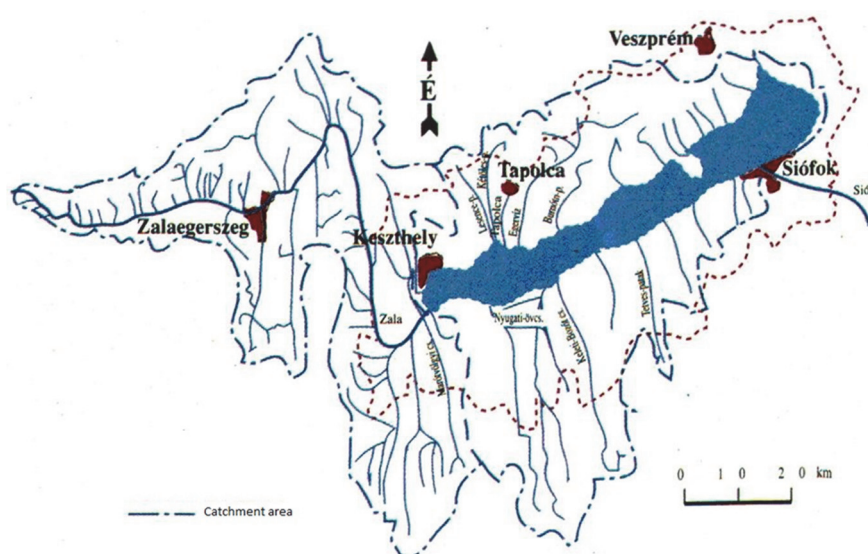


Fig. 1. Catchment area of the Balaton Lake.

(source: website of the Hungarian Water Inspectorate)

Several examples can be found in the literature for application of the Mann-Kendall trend test, e.g., *Patle and Libang* (2014) argued on trend analysis of annual and seasonal rainfall the northeastern region of India, *Salmi et al.* (2002) analyzed the trends of atmospheric pollutants in Finland. Meteorological applications can be read in *Rahman and Begum* (2013), who determined trends of rainfall of the largest island in Bangladesh. *Ganguly et al.* (2015) investigated the tendencies of rainfall in Himachal Pradesh (state in North India) between 1950 and 2005. *Gavrilov et al.* (2015a, 2015b, 2016) examined trends of air temperature by Mann-Kendall test in Vojvodina, Serbia. *Salami et al.* (2014) applied this non-parametric trend test for the analysis of hydro-meteorological variables in Nigeria. *Mapurisa and Chikodzi* (2014) made an assessment of trends of monthly and seasonal rainfall in Southeast Zimbabwe. *Karmeshu* (2012) investigated the temperature- and precipitation changes in the northeastern part of United States. Hydrological utilization is given by *Hamed* (2008). *Burn and Hag Elnur* (2002) estimated the trends and variability of 18 hydrological variables by Mann-Kendall trend test. *Hirsch et al.* (1991) used the method for investigation of stream water quality.

2. Data and methods

Monthly amounts of precipitation were analyzed from 1871 to 2014 measured at the beginning in the territory of the ancient Georgikon Academy of Agriculture at Keszthely, then at the meteorological station of the Hungarian Meteorological Service. The dataset was provided by the Department of Meteorology and Water Management of University of Pannonia Georgikon Faculty (Keszthely). This dataset is special, because few stations in Hungary have continuous measurements of more than 140 years with detailed historical background (*Kocsis and Anda*, 2006). The dataset was analyzed in three sessions: annual amounts, seasonal amounts and monthly precipitation sums. Seasons of temperate climate were performed as common in meteorology, e.g., spring: March, April, May, etc.

2.1. The Mann-Kendall trend test

Mann-Kendall trend test is widespread in climatological and hydrological time series analysis, because it is simple and robust, it can cope with missing values and values under detection limit (*Gavrilov et al.*, 2016). This non-parametric test is commonly used to detect monotonic tendencies in series of environmental data, too (*Pohlert*, 2016). This method has no requirement for the distribution of the population, as the regression method has a requirement for normal distribution. No assumption of the normality is required (*Helsel and Hirsh*, 2002). *Hirsch et al.* (1982) and *Hirsch and Slack* (1984) developed the method and introduced seasonal Mann-Kendall test for data that are serially dependent. *Hamed and Rao*

(1998) developed a modified Mann-Kendall test for autocorrelated data. *Yue et al.* (2002) investigated the power of the Mann-Kendall test in hydrological series.

The Mann-Kendall trend test is based upon the work of *Mann* (1945) and *Kendall* (1975), and is closely related to the Kendall's rank correlation coefficient. The methodology is introduced following the detailed descriptions given by *Gilbert* (1987) and *Hipel and McLeod* (1994) as follows:

In case of determining the presence of monotonic trend in a time series, the null hypothesis (H_0) of the Mann-Kendall test is that the data come from a population where random variables are independent and identically distributed. The alternative hypothesis (H_a) is that the data follow a monotonic trend over time. The Mann-Kendall test statistic is given as

$$S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sign}(x_j - x_k), \quad (1)$$

where $j > k$, and

$$\text{sign}(x) = \begin{cases} +1, & x > 0 \\ 0, & x = 0 \\ -1, & x < 0 \end{cases}. \quad (2)$$

Kendall (1975) proved that S is asymptotically normally distributed with the following parameters (mean and variance):

$$E[S] = 0$$

$$\text{Var}[S] = \{n(n-1)(2n+5) - \sum_{j=1}^p tj(tj-1)(2tj+5)\}/18, \quad (3)$$

where p is the number of the tied groups in the data set, tj is the number of data in the j th tied group, and n is the number of data in the time series.

Positive value of S means that there is an increasing trend, negative value of S means the opposite, there is a decreasing trend with time. It was proven that if the number of data is above 10, the standard normal variate Z can be used for hypothesis test:

$$Z = \begin{cases} \frac{S-1}{[\text{Var}(S)]^{1/2}}, & \text{if } S > 0 \\ 0, & \text{if } S = 0 \\ \frac{S+1}{[\text{Var}(S)]^{1/2}}, & \text{if } S < 0 \end{cases}. \quad (4)$$

During the hypothesis test, $\alpha=5\%$ significance level was used in the two-tailed and one-tailed tests as well, and statistical significance was determined. S is closely related to the Kendall's rank correlation coefficient (τ):

$$\tau = \frac{S}{D}, \quad (5)$$

where D is the possible number of data pairs from n members of the dataset:

$$D = \binom{n}{2}. \quad (6)$$

After detecting the non-parametric trend, the Sen's slope estimator was applied. This is a non-parametric method that can calculate the change per time unit (direction and volume). Detailed description of the method is given by *Sen* (1968). Not only the value of the slope, but also the confidence interval was estimated under $1-\alpha=95\%$ probability level.

SPSS 17.0 and Addinsoft's *XLSTAT* (2016) was used for carrying out the computations.

3. Results

Our previous studies focused on linear trend approximation for analyzing the long-time series of precipitation measured at Keszthely station (Hungary) (*Kocsis*, 2008; *Kocsis* and *Anda*, 2016). Results are summarized in *Table 1*. It can be stated that statistically significant changes could be detected by linear regression method with parametric test for the slope by two-tailed t-test for spring, and among the monthly data of April and October. *Table 2* presents that the requirements for the normal distribution for the population are fulfilled, tested by the Kolmogorov-Smirnov test for distribution. However, in some cases (February, March, July, and November), the distribution cannot be accepted as normal distribution according to p-value ($\alpha=5\%$). For this reason, non-parametric methods were used to analyze all the data set (annual sums, seasonal and monthly sums).

Table 1. Summary of the previous findings using linear trend and two tailed t-test for testing the significance of the slope

	Significant tendency in 100 years		
	1871–2000 (Kocsis, 2008)	1871–2010 (Kocsis, 2015)	1871–2014 (Kocsis and Anda, 2016)
Changes in annual sum	none	none	none
Changes in seasonal sums	Spring (–35 mm)	Spring (–29 mm)	Spring (–31.5 mm)
Changes in monthly sums	October (–26 mm)	October (–25 mm)	April (–14 mm) and October (–23.5 mm)

Table 2. Results of the test of normal distribution

Distribution (Kolmogorov-Smirnov test)		
H0: The sample follows a normal distribution		
<i>Ha: The sample does not follow a normal distribution</i>		
	Hypothesis accepted at $\alpha=5\%$	p-value
Annual sum	H0	35.2%
Spring	H0	42.1%
Summer	H0	80.0%
Autumn	H0	53.1%
Winter	H0	91.0%
January	H0	34.7%
February	<i>Ha</i>	4.7%
March	<i>Ha</i>	3.6%
April	H0	5.3%
May	H0	28.6%
June	H0	75.2%
July	<i>Ha</i>	2.5%
August	H0	25.6%
September	H0	47.4%
October	H0	5.7%
November	<i>Ha</i>	3.1%
December	H0	7.1%

As the first step, the Kendall's tau rank correlation coefficient was determined to show the sign of the relationship between time and the variables. Tau was first tested by the two-tailed test (*Table 3*). Significance level was stated at $\alpha=5\%$. In two-tailed tests, p-values proved significant negative relationships for spring, autumn, and among the monthly sums, for October. P-values of some other monthly sums were near to the limit, so one-tailed tests according to the sign of Kendall's tau were also used. P-values in one-tailed tests show the presence of negative relationship in the annual sum, precipitation sums in spring, autumn, April, May, and October, and positive relationship for the winter (*Table 3*).

Table 3. Values of Kendall's rank correlation coefficient and their significance

	Kendall τ	Two-tailed test ($\tau \neq 0$)	One-tailed test ($\tau < 0$ or $\tau > 0$)
		p-value	p-value
		*significant at $\alpha=5\%$	
Annual sum	-0.096	8.9%	4.5%*
Spring	-0.159	0.5%*	0.2%*
Summer	-0.036	51.9%	26.0%
Autumn	-0.112	4.6%*	2.3%*
Winter	0.096	8.8%	4.4%*
January	0.042	45.2%	22.6%
February	0.063	26.1%	13.0%
March	-0.054	33.9%	16.9%
April	-0.107	5.7%	2.8%*
May	-0.104	6.4%	3.2%*
June	0.01	86.3%	43.1%
July	-0.018	74.4%	37.2%
August	-0.042	45.3%	22.7%
September	-0.054	33.4%	16.7%
October	-0.153	0.7%*	0.3%*
November	0.012	83.4%	41.7%
December	0.061	27.7%	13.9%

Based on the Kendall's tau (τ), the Mann-Kendall trend test was used to detect monotonic trends in the time series. First two-tailed (*Table 4*) and then one-tailed z-tests were applied according to the sign of the calculated S value, respectively (*Table 5*). Decision about the statistical significance was made using

empirical significance level (p-value) compared to $\alpha=5\%$. The two-tailed Mann-Kendall test (with the alternative hypothesis that there is no trend in the time series) gave significant result for spring, autumn, and October in concordance with the results for the tests for τ . Seasonally in spring and autumn, while among the months, in October the presence of significant monotonic trend could be found. One-tailed tests have an alternative hypothesis, that in the time series, a negative/positive trend is present according to the sign of computed S statistics, respectively. These tests gave the results in accordance to the one-tailed tests for τ , as S and τ statistics are related. According to p-values, significant declining trend can be experienced in the annual sum, precipitation sums in spring, autumn, April, May, and October, and increasing trend can be shown for the winter.

Table 4. Results of the two-tailed Mann-Kendall trend test

Mann–Kendall trend test (two-tailed)		
H0: There is no trend in the series		
<i>Ha: There is a trend in the series</i>		
	Hypothesis accepted at $\alpha=5\%$	p-value
Annual sum	H0	9.0%
Spring	<i>Ha</i>	0.5%
Summer	H0	52.1%
Autumn	<i>Ha</i>	4.6%
Winter	H0	8.8%
January	H0	45.3%
February	H0	26.1%
March	H0	33.9%
April	H0	5.7%
May	H0	6.4%
June	H0	86.4%
July	H0	74.5%
August	H0	45.4%
September	H0	33.5%
October	<i>Ha</i>	0.7%
November	H0	83.6%
December	H0	27.8%

Table 5. Results of the one-tailed Mann-Kendall trend test

Mann-Kendall trend test (one-tailed)			
H0: There is no trend in the series			
<i>Ha: There is a positive or negative trend according to the sign of S</i>			
	Hypothesis accepted at $\alpha=5\%$	p-value	S
Annual sum	<i>Ha</i>	4.5%	−983
Spring	<i>Ha</i>	0.2%	−1633
Summer	H0	26.0%	−373
Autumn	<i>Ha</i>	2.3%	−1154
Winter	<i>Ha</i>	4.4%	988
January	H0	22.7%	435
February	H0	13.1%	651
March	H0	17.0%	−554
April	<i>Ha</i>	2.8%	−1103
May	<i>Ha</i>	3.2%	−1073
June	H0	43.2%	100
July	H0	37.3%	−189
August	H0	22.7%	−434
September	H0	16.8%	−559
October	<i>Ha</i>	0.3%	−1568
November	H0	41.8%	121
December	H0	13.9%	629

The Sen's slope estimator was used to determine the slope of the trend (linear change per time unit). Confidence interval is also given for the slope. This interval contains the real value of the slope of the tendency in the population by the probability of 95% (Table 6). According to the one-tailed Mann-Kendall test, the decreasing tendencies are the follows for 100 years on average: the annual sum decreased by 44 mm, the precipitation sum of spring and autumn declined by 29.7 mm and 25.7 mm, respectively. The rainfall sum of winter increased by 15.4 mm. Sums of April, May, and October declined by 10.8 mm, 13 mm, and 20.9 mm, respectively.

Table 6. Values of the Sen's slope and their confidence interval

	Sen' slope	Lower boundary of the confidence interval at $1-\alpha=95\%$ probability level	Upper
Annual sum	-0.444	-0.703	-0.207
Spring	-0.3	-0.396	-0.2
Summer	-0.095	-0.231	0.046
Autumn	-0.26	-0.38	-0.148
Winter	0.156	0.075	0.23
January	0.028	0	0.063
February	0.047	0.008	0.089
March	-0.043	-0.082	0
April	-0.109	-0.163	-0.059
May	-0.131	-0.192	-0.067
June	0.013	-0.063	0.08
July	-0.022	-0.098	0.036
August	-0.068	-0.139	0.011
September	-0.074	-0.13	-0.002
October	-0.211	-0.274	-0.144
November	0.012	-0.041	0.073
December	0.05	0.005	0.095

4. Discussion

Parametric methods (linear trend, t-test for slope) for analyzing time series are the simplest methods to get insight to the changes of a variable over time. These methods have a requirement for normal distribution of the population that can be a limit for application. Non-parametric methods are distribution-free methods, and researchers can have a more sophisticated view to the changing tendencies. Our previous results can be accepted, but the Mann-Kendall trend test proved other signs of precipitation changes at the examined meteorological station. One-tailed Mann-Kendall trend tests gave the evidence that decreasing tendencies can be statistically proven in the annual sum by 44 mm for 100 years on average, in the precipitation sum of spring and autumn by 29.7 mm and 25.7 mm, respectively. The rainfall sum of winter increased by 15.4 mm. Sums of April, May, and

October declined by 10.8 mm, 13 mm, and 20.9 mm, respectively. These results are parallel to the literature in the direction of the precipitation changes, but are not clearly in accordance regarding the volume of the changes estimated for the region of the examined meteorological station (Keszthely, Trans-Danubia, Hungary). Results of two-tailed tests of previous parametric and recent non-parametric methods are easily comparable. Parametric method (linear trend) proved significant tendencies for spring, April, and October (decreasing tendencies). Two-tailed non-parametric Mann-Kendall trend tests show significant tendencies for spring, autumn, and October (declining trends). Advantage of the non-parametric methods, that is fewer requirements for application, and the results derived from this research are more consistent with the literature. Declining tendencies of precipitation sums in spring and autumn possibly have unfavorable effect on the water-budget of the soil and Balaton Lake.

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